Aerosol Remote Sensing and Atmospheric Correction Beyond VIIRS: What will Expanded Wavelengths, Hyperspectral, Multi-angle, and Polarization Bring to the Table

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Global Aerosol Optical Depth at 550 nm from VIIRS (IDPS) 13 August 2017



NOAA Global View

Global Aerosol Optical Depth at 550 nm from Aqua MODIS (DT) 13 August 2017



NASA World View



Global Aerosol Optical Depth at 550 nm from Aqua MODIS (DT) 2016 Annual Mean



NASA Giovanni



 $\begin{array}{l} \mu_{o} = \cos(\text{solar zenith angle}) \\ \mu = \cos(\text{sensor zenith angle}) \\ \phi = \text{relative azimuth} \\ \tau^{*} = \text{aerosol optical thickness} \end{array} \qquad \begin{array}{l} R = \text{surface reflectance} \\ E_{o} = \text{extraterrestrial} \\ \text{irradiance} \\ T_{t} = \text{total transmission} \\ S = \text{spherical albedo} \end{array}$



 $\mu_o = \cos(\text{solar zenith angle})$ $\mu = \cos(\text{sensor zenith angle})$ $\phi = \text{relative azimuth}$ $\tau^* = \text{aerosol optical thickness}$

R = surface reflectance

 $T_{\rm t}$ = total transmission

S = spherical albedo

This is a single channel Dark Target retrieval. (one wavelength at a time)

Requires assumptions of particle size and refractive index, as well as surface reflectance

Returns 1 piece of information: τ^*





AOD 550 nm Aqua MODIS Monthly mean July 2017

Fine mode only

Total AOD

MODIS/Aqua MYD08_M3.A2017182.006.2017213173821.hdf

none

Global Aerosol Optical Depth at 483 nm from Aura OMI 2016 Annual Mean





Radiative transfer in the near UV



$$\Omega$$
 = ozone content
p_o = surface pressure

Note: no dependence on aerosol in the equation

Radiative transfer in the near UV

$$L^{*}(\Omega,\mu,\mu_{0},\varphi,p_{0}) = L^{0}(\Omega,\mu,\mu_{0},\varphi,p_{0}) + \frac{RT_{t}(\Omega,\mu,\mu_{0},p_{0})}{1-RS(\Omega,p_{0})}$$
$$R_{\lambda_{0}} = \frac{L^{*}_{\lambda_{0}} - L^{0}_{\lambda_{0}}}{T_{t}_{-\lambda_{0}} + S_{\lambda_{0}}(L^{*}_{\lambda_{0}} - L^{0}_{\lambda_{0}})}$$

Reflectance of the lower surface in one wavelength.

If no aerosol, and ozone absorption is minimal (330-390 nm), then just Rayleigh scattering. L⁰, T_t and S are obtained from Rayleigh scattering calculations

$$R_{\lambda_{0}} = \frac{L_{\lambda_{0}}^{*} - L_{\lambda_{0}}^{0}}{T_{t_{-}\lambda_{0}} + S_{\lambda_{0}}(L_{\lambda_{0}}^{*} - L_{\lambda_{0}}^{0})}$$

Assume lower surface reflectance has NO spectral dependence.

Calculate R at two wavelengths.

All spectral dependence should match Rayleigh expectations. If not then, **AEROSOL**.

Another SCIAMACHY scene of smoke over broken clouds over ocean



Measurements deviate from Rayleigh-only expectations, In <u>magnitude</u> and in <u>spectral dependence</u>.

Global Aerosol Single Scattering albedo at 483 nm from Aura OMI 2016 Annual Mean





One additional problem with the UV approach...

In the UV, deviation from Rayleigh depends on the height of the aerosol.

2 pieces of information/ 3 important parameters:

- Loading (AOD)
- Absorption (SSA)
- Height (h)

The dark target VNIR-SWIR approach...

- Quantify the deviation from background radiance
- Spectral dependence gives size.

The UV approach...

- Quantify the deviation from Rayeigh scattering
- Spectral dependence gives absorption

Combining approaches (MODIS + OMI) we should be able to retrieve **all 3 parameters: AOD, SSA and h**

Retrieval of aerosol height (combined L2 products, not radiances)



Satheesh, Torres, Remer et al., JGR, 2009



Plankton, Aerosol Cloud ocean Ecosystems NASA mission for 2022





PACE will show all chlorophyll is not created equal

OCI Specifics:

- Single detector, rotating telescope scanner (like SeaWiFS)
- 20-degree tilt to avoid sun glint
- Monthly lunar calibration of all science detectors
- Ground sample distance ~ 1 square kilometer at nadir
 - 5 nanometer (nm) resolution from 350 to 890 nm
 - Plus short-wave infrared (SWIR) bands
 centered on 940, 1240, 1380, 1640, 2130 &
 2250 nm

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1 km

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Mattoo and Remer (in progress)

Expected for PACE from OCI

Spectral AOD, SSA and height over oceans (when AOD_550 is sufficiently high ~0.30)

Land not yet explored.

In forward processing, at 3 to 10 km, Across the OCI swath (2-day coverage)

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Using Oxygen A-band to retrieve aerosol height

DOAS = differential optical absorption spectroscopy

O₂: both the amount and cross-section of the gas are accurately known.

DOAS then tells us about the height of the scattering layer (0-2 km vs. 3-5 km)

Davis, Kalashnikova, in progress

Plankton, Aerosol Cloud ocean Ecosystems NASA mission for 2022



<u>Primary instrument: OCI for ocean color,</u> but also an excellent aerosol instrument:

- Everything for aerosol that MODIS and VIIRS offers
- Except the thermal channels for cloud mask
- Plus, UV channels at 1 km
- And Oxygen A-band at 5 nm

New opportunities for aerosol absorption and height

Secondary instrument: Multi-angle imaging polarimeter



Figure 6. Left. Calculated degrees of freedom information content) from **From K. Knobelspiesse** andidate multi-angle polarmeter, the

Theoretical calculations Theoretical calculations Present different chlorophyll OCI (single angle radiometer) Sottin black duncertainties in Sottin black duncertainty Sott





UMBC PI-Neph







Espinosa, Dolgos and the UMBC team

UMBC PI-Neph measurements of 3 different fine mode aerosols from SEAC4RS



Espinosa et al., 2017

Volume size distribution



GRASP retrieval applied to PI-Neph measurements

Espinosa et al. (to be submitted)

GRASP: Generalized Retrieval of Aerosol and Surface Properties



Aerosol retrievals from airborne multi-angle polarimeter (AirMSPI)



AirMSPI on NASA ER-2 over Fresno CA (in red)

Near-time AERONET retrievals at Fresno station (other colors)

Xu et al. (2017)



Xu et al. 2017

Where are we heading?

Towards broader aerosol-relevant spectra; Finer spectral resolution; Finer spatial resolution for UV; Multi-angle polarimeters; Retrievals of more aerosol properties; With better accuracy; Expanding retrievals to over clouds and difficult surfaces; Simultaneous retrievals of aerosol + surface; Advanced retrieval algorithms that can handle all this information

How will we get there?

- PACE OCI: 2022, but in political limbo
- PACE polarimeter: even less assured
- MetOP 3MI: 2018 and continuing
- MAIA: an instrument without a launch date
- HARP Cubesat: early 2018

Photos of Actual Instrument & Spacecraft



Preliminary Results from LMOS, <u>12 June 2017</u> – AirHARP Zion



Example of Multiangle observation

Full overlap region





